Massachusetts Coastal Zone Management

MASSACHUSETTS COASTAL SUBMERGENCE PROGRAM

Passive Retreat of Massachusetts Coastal Upland
Due to Relative Sea-Level Rise
PASSIVE RETREAT OF MASSACHUSETTS COASTAL UPLAND DUE TO RELATIVE SEA-LEVEL RISE

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INTRODUCTION

Shoreline recession is recognized widely as a major environmental management issue in Massachusetts as well as in many other parts of the United States and throughout the world (Bird, 1976). In considering this issue, it is essential to separate the retreat of coastal upland areas from the retreat of wetlands because of the differences between the processes involved. The retreat of a barrier beach, for example, may involve the landward translation of an entire feature without diminution in its size, but upland retreat always results in the loss of upland area. Although upland loss usually is accompanied by wetland gain, the upland lost is an irreversible loss of that area from those land uses for which wetlands are considered unfit. In Massachusetts these uses include, for example, human habitation, transportation and commerce.

Coastal upland retreat takes two distinct forms: active wave-produced erosion and passive loss resulting from relative sea-level rise. While a rise in relative sea level contributes to active wave-produced erosion, it is not possible at present to quantify the contribution to erosion made by sea-level rise. On the other hand, the recession of a passive shoreline as sea level rises can be estimated with reasonable accuracy.

Unfortunately, estimates of passive shoreline recession are seldom available, probably because upland loss due to this cause generally is considered to be small compared to that due to erosion. Relative sea-level rise along the Massachusetts coast over the past 40 years ranges between 2 and 3 mm. per year (Aubrey and Emery, 1983). Within recent years, however, a rapidly increasing body of data has appeared in support of the hypothesis that global climatic warming within the next century will cause increasing global sea level rises that can not be ignored. Hoffman (1984), for example, has projected global sea-level rises by the year 2100 ranging from 1.8 ft. ("low scenario") to 11.3 ft. ("high scenario").

Some emphasis in this report is placed on relative sea-level rise rather than absolute sea-level rise. Coastal submergence results not only from rise of ocean levels, but also from sinking of the land. In Massachusetts, nearly two-thirds of the submergence during the past century (documented by tide-gauge data) results from subsidence of the land. Only one-third of the submergence appears
considerable uncertainty in the scientific basis for predicting the details of global warming, how can these uncertainties be translated into an equitable planning or zoning process? That a global warming is in process and will continue is incontrovertible. What are not known precisely are the magnitude and timing of this global warming, and its exact impact on sea levels.

The appropriate response to these issues and results on local and state-wide levels is one of increasing awareness. Legislation and re-zoning may be premature. However, awareness by town planners, politicians, and Conservation Commissions, for instance, must be increased. Long-range planning could take these shoreline retreat data into account when making major land use decisions. Conservation Commissions could err on the side of caution in a coastal construction issue, mandating pile foundations in areas of critical concern. Public works could incorporate these data in siting wells or new sewer systems. In summary, some rational response to these sea-level rise issues are appropriate at this time. Major legislation and drastic changes in regulations, however, may be premature and might better await a clearer consensus from the scientific community before enactment.

Users of the data presented in this report must be aware that passive shoreline retreat via inundation is not the sole effect arising from global warming to which coastal communities must respond. Although the present study considers only the effect of passive retreat due to inundation, other impacts may be equally important. For example, rising sea levels will change the base level for river drainage and groundwater flow. Water quality deterioration may result from this impact. In addition, global warming will raise the ocean surface temperature, increasing the size of the "warm-pool" of water that is responsible for generating tropical cyclones. Although difficult to predict in detail because of the complexities of non-linear atmospheric physics, this ocean warming is certain to alter storm climates along the eastern seaboard and elsewhere. If the net product is an increase in tropical cyclones reaching the northeast, this could result in more severe short-term (order of decades) economic impact than that due to simple passive retreat. While the present study investigates an issue of fundamental importance, the user should be aware of these other significant impacts, and plan their rational response to global warming accordingly.

**METHODS**

Quantification of the passive retreat of coastal upland presents special problems due to the peculiar "fractal" nature of the passive shoreline (Mandelbrot, 1977). Simply stated, the problem is that the complex form of the passive shoreline does not simplify as smaller and smaller segments are examined, and thus the "tangible" shoreline always remains just out of reach of the investigator who would measure it. In order to skirt this problem, the present study deals not with the linear retreat of the shoreline, but rather with the areas that are lost as the shoreline recedes. Two separate approaches are used, each having special advantages and disadvantages. In the first, which treats entire coastal communities, the distribution of the area of the community with respect to its elevation is presented in the form of "hypsometric" curves, or cumulative frequency diagrams.
While this is a powerful tool for the analysis of such geographical units as a whole, the results give no information about the change at a specific point within that unit. The second approach makes use of color-coded maps of areas that are of special concern for the management of ports and harbors. For this purpose the harbors of Hyannis, Westport and Gloucester were chosen. While it is difficult to quantify the effects of small changes from these color-coded maps, the areas that will (and will not) be affected are displayed clearly.

**Hypsometry**

As a tool for calculating the retreat of coastal upland resulting from relative sea-level rise, hypsometry has been discussed by Giese et al. (1985). Unlike previous work, however, the present study makes use of digital elevation data that permits the application of the hypsometric method to large areas. A separate hypsometric curve was calculated for each of 72 Massachusetts coastal communities.

The initial data for the upland hypsometric calculations were obtained from the U.S. Geological Survey's (USGS) National Cartographic Information Center (NCIC). They consist of two separate types of digital information, both of which are stored on magnetic tapes. The first type is elevation data that consists of land surface elevations to the nearest meter arranged in south-to-north profiles for entire one-degree latitude by one-degree longitude areas. The data points within the profiles, as well as the profiles themselves, are separated by intervals of three arc seconds, which is equivalent to a distance of about 92 m in a north-south direction and about 69 m in an east-west direction at a latitude of 42 degrees (the approximate mid-point of the study area).

The second type of digital data consists of land use and land cover codes arranged in west-to-east rows aligned along a Universal Transverse Mercator (UTM) grid and covering entire one-degree latitude by two-degree longitude areas. The UTM coordinate system is rotated slightly counterclockwise with respect to the geographic latitude-longitude coordinate system. The land use and land cover code data points, and the rows containing them, are separated by intervals of 200 m. The land use codes include the U.S. Bureau of the Census designation for each 200 m square; these data permit the assignment of each square to a specific town or city. The land cover classification codes are sufficient to permit exclusion of wetland and inland water areas.

A large part of the effort for this study consisted of the programming required to combine the raw digital data described above to produce a single data set consisting of elevation, census code and land-cover code for each 3-second box within one-degree blocks. A description of the programs and their use is included in Appendix B.

During the study, the accuracy of the census and land-cover codes was checked by reference to the appropriate U.S.G.S. 7.5-minute series topographic maps, as well as by comparing the total calculated upland area of individual communities with the known value of their total land area. No problems were encountered with either type of code. Unfortunately, the same was not true of the elevation data. Initial tests of these data were performed by comparing profiles derived from the digital data to profiles based on the 7.5-minute series maps. The results of these tests generally
were satisfactory, particularly considering the fact that the entire elevation data set for each community was to be combined. However, when the cumulative distributions of elevation data were completed, it was evident that the USGS data were biased toward maxima in the vicinity of 3, 15, 30, 45 and higher multiples of 15 m. A program was written to smooth the distributions by redistributing the excessive values linearly to the depleted elevation categories between the maxima. A description of this procedure is included in Appendix B. The hypsometric curve thus calculated for one community (Barnstable) then was compared to the curve derived by a graphical method, and found to be acceptable. Nevertheless, it must be noted that cumulative hypsometric data presented in this report are less accurate than those that could be obtained using unbiased elevation data.

**Color-coded Maps**

The three maps that accompany this report were prepared to illustrate the effect upon three harbors of the relative sea-level rise predicted by four different scenarios for the year 2100. The harbors of Hyannis, Westport and Gloucester were chosen for this purpose because of their contrasting geological settings and because of their distribution along the Massachusetts coast.

These maps were generated using data derived from the digitization of selected portions of the 7.5-minute series topographic maps for the three harbors. These maps have a contour interval of 10 feet, which is too great to resolve the flooding that was to be shown. Therefore, a surface was modelled to fit the digitized contours using a modified form of existing software. The levels of flooding characterizing the four scenarios then were applied to this modelled surface. Using color-plotting software and equipment, the flooded areas were displayed on color-coded maps. A detailed description of the methodology employed is included in Appendix B.

The four sea-level rise scenarios illustrated on the maps were presented by Hoffman (1984), and produce flooding of 1.8, 4.7, 7.1 and 11.3 feet by the year 2100. These values were added to the NGVD elevations of the mean high water shorelines shown on the 7.5 minute series maps. The shoreline elevations were assumed equal to the half-tidal range at each particular harbor plus 0.5 ft to account for relative sea-level rise since 1929, the date of the NGVD datum. Local variations between NGVD and mean sea level were ignored, although these data are available.

Two important differences between the hypsometric calculations and the color-coded maps should be noted. First, while the hypsometric calculations refer only to coastal uplands and wetland areas are entirely excluded, the color maps use as their basic reference level the present mean high water shoreline that, in many areas, borders on coastal wetlands. Therefore, the areas shown as being flooded according to the lowest rise scenario include wetland areas, many of which are salt marshes. The second difference, discussed in more detail below, is that the maps include a consideration of the ground water table rise that accompanies a rising relative sea level. This effect is excluded from the calculations based on hypsometry.
Figure 1: Schematic of datum planes selected for sea-level rise scenarios.
RESULTS AND DISCUSSION

The hypsometric curves for each community, together with tables giving the cumulative distribution of upland area with respect to elevation for each, are presented in Appendix A. The first area value presented in each table and graph (that for 3 m) represents the upland area that lies between 2.5 and 3.5 m. This interval was chosen because at lower elevations it is impossible to distinguish between upland and wetland in the source data, and does not imply that there is no upland below 2.5 m. in the community. The assumption is made throughout this study that the areal frequency of upland below 2.5 m. is equal to that at 3 m. No assumption is made, however, about the elevations of the wetland/upland boundaries within the community, other than that these boundaries, whatever their elevations, rise at the same rate as relative sea level (figure 1). It also should be noted that the data terminate at an elevation of 60 m., even when higher land exists within a community, in order to limit the size of the figures.

There is a striking variation between communities in the shape of their hypsometric curves, reflecting variation in the geological processes that formed them. For example, communities on glacial outwash plains, such as Yarmouth, have curves with flatter slopes at low elevations as compared to those, such as Brewster, that lie on glacial moraines. Certain well-known local topographic features, such as the "Wellfleet Plains", also show up clearly on the figures.

Making use of these hypsometric data, calculations have been made of the upland areas that each community would lose given particular changes in relative sea level. The results of these calculations are presented in Table 1. The first column in Table 1 lists the names of the coastal communities of Massachusetts, and the second column gives the upland area, in acres, of each community. The third column lists the percentage of upland area - and the fourth column the actual area measured in acres - that each community loses in response to a relative sea-level rise of 0.01 ft. (3 mm), considered here to be the historical mean annual rate of rise (Aubrey and Emery, 1983). The following three pairs of columns give the amount of retreat, first in percent of total upland area and then in acres, that will occur between 1980 and 2025 given three different sea-level rise scenarios. The first scenario, case 1, calls for a continuation of the historical mean annual relative sea-level rise rate of 0.01 ft/yr, giving a total rise of 0.45 ft over the 45 year period. Case 2 assumes that global sea level will rise 0.86 ft over the 45 year period (as given by Hoffman's "mid-range low" scenario) and that the local coastal subsidence rate will remain at 0.0062 ft/yr, giving a total relative rise of 1.14 ft by 2025. Case 3 is based on the same assumption about local subsidence, but uses Hoffman's "mid-range high" global sea-level rise estimate of 1.29 ft by 2025, yielding a total relative rise of 1.57 ft.

The total Massachusetts upland loss at the historical relative sea-level rise rate is 65.4 acres per year. Averaged among the 72 communities, this works out to be 0.9 acres per year per community. However, the variation between communities is great, covering two orders of magnitude: Nantucket loses 6.1 acres per year, while Winthrop loses only 0.06 acres. After Nantucket, other communities having large annual losses are: Wareham, 4.7 acres; Falmouth, 3.8
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TABLE 1 (continued)

CALCULATED UPLAND RETREAT
(Areas are in acres, % represents percent of upland submerged)

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<tr>
<th>TOWN NAME</th>
<th>AREA (ACRES)</th>
<th>% AREA</th>
<th>AREA 0.01 ft/yr RISE</th>
<th>% AREA</th>
<th>AREA 0.45 ft RISE</th>
<th>% AREA</th>
<th>AREA 1.14 ft RISE</th>
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The following coastal towns loose less than 0.001% of their total upland area annually as the result the historical mean sea-level rise rate of 0.01 ft/yr, and therefore were omitted from this table: Braintree, Hanover, Milton, Norwell, Peabody and Pembroke.
acres; Barnstable, 3.7 acres; and Yarmouth, 3.2 acres. In terms of annual percentage of total upland lost per year, the communities most affected are: Marion, which loses 0.031% per year, followed by Nantucket which loses 0.027% per year, and Hull and Yarmouth, which lose 0.026% per year.

Looking forward to the year 2025, if the historical rate of relative sea-level rise were to remain unchanged (case 1), the total Massachusetts upland loss would be 2,945 acres. A relative sea-level rise of 1.14 ft, as projected in case 2, would be accompanied by an upland loss of 7,459 acres, and a rise of 1.57 ft (case 3) would cost the commonwealth 10,273 acres of upland.

When considering these figures, it is important to realize that they do not include the upland losses that would result from the response of ground water levels to sea-level rise. In those communities where bedrock is absent and the terrain consists of unconsolidated sediments, the water table level over geological time periods is controlled by relative sea level. As sea level rises, the water table level rises with it, increasing the size of existing streams, ponds and bogs, and creating new ones. This effect has not been included in the hypsometric analysis discussed above, although it was taken into account in the construction of the color-coded maps.

The reader also should bear in mind that the calculated upland retreat rates are based on the assumption that the coastal uplands have a natural form and are not protected by engineering structures. Particularly in urban coastal areas where seawalls, riprap and fill are prevalent, the actual losses will be less than those predicted here. As the color-coded maps indicate, however, when large values of sea-level rise are considered, these structures are overwhelmed.

It is of interest that the presently existing rate of upland retreat due to the passive effects of relative sea-level rise is much greater than the upland retreat rate due to active wave-produced erosion. This may be illustrated by a consideration of the Cape Cod coast, which is well-known as a region of rapid erosion. While detailed estimates for cliff retreat do not exist for the entire region, the rate of erosion of the outer coast is well-known (e.g., Zeigler et al., 1964), and reasonable estimates can be made for the remaining and more slowly retreating cliff areas. Using such existing information and reasonable estimates, the annual upland loss experienced by Cape Cod as the result of active wave-produced erosion is about 9 acres per year. On the other hand, the annual loss due to the passive effects of relative sea-level rise, calculated from the figures for each Cape Cod town listed in Table 1, is about 24 acres per year. Thus it is seen that even considering a region of rapid erosion, and excluding the effects ground water table rise, passive retreat accounts for 73% of coastal upland loss under present conditions.

Figures 2, 3, and 4 present the color-coded maps depicting the submergence patterns of Hyannis, Gloucester, and Westport harbors that would accompany each of the four Hoffman (1984) sea-level rise scenarios for the year 2100. The maps show in red the land areas that would be lost given the low scenario rise of 1.8 ft, in yellow the submerged areas given the mid-range low scenario rise of 4.7 ft, and in green and blue the areas submerged by the mid-range high scenario rise of 7.1 ft and the high range scenario rise of 11.3 ft respectively. The low scenario changes are
extensive only in wetland areas, such as the salt marshes northwest of Gloucester Harbor, the sand spit southwest of Hyannis Harbor, and fringing marshes in Westport Harbor. While the upland lost given this scenario is not extensive, the increased potential for storm wave and flooding damage should be of concern.

The submergence that would accompany the other scenarios is extensive and would impact severely operations of harbor facilities. In addition, the maps show locally significant flooding of inland areas for these scenarios resulting from elevated ground water levels. As has been discussed above, it should be kept in mind that the levels used in applying these scenarios do not include the effects of coastal subsidence, and that for the lower rise rates the increases would be significant were they to be included.

CONCLUSIONS

Major conclusions of the present study are:

1. Relative sea-level rise is the major process responsible for upland loss in Massachusetts. Neglecting coastal erosion and fresh water table changes, Massachusetts presently looses about 65 acres of upland each year due to passive submergence.
2. The rate of upland loss due to passive submergence varies widely from town to town, and depends upon the geology of the region in which the town lies.
3. The hypsometric curves of the towns provide important basic information that permits the calculation of the upland areas which those towns will lose to passive submergence as the result of any given increase in relative sea-level.
4. The total land loss by the year 2025 has been calculated for several relative sea-level rise scenarios. At the present rate of rise, Massachusetts will have lost about 3,000 acres of upland between 1980 and 2025. This is the same upland loss that occurred between 1935 and 1980, an equal length of time. For a rise of 1.14 ft, about 7,500 acres would be lost; and for a rise of 1.57 ft, the maximum likely, over 10,000 acres would be lost. Given a nominal value of ocean-front property of $1,000,000 per acre, the economic impact of this retreat is substantial.
5. Color-coded maps are a useful device for depicting the specific areas that will be submerged as the result of specified increases in relative sea level. These maps could be developed for each coastal town in the future, to provide guidance for land use, public works, and conservation decisions.
6. These data can be used immediately to help provide a rational basis for local response to global climate warming. Data from this report, although representing hypothetical scenarios, remove the quantification of the impacts of passive retreat from the realm of speculation, placing them on a firmer basis. Although enactment of legislation and major revision of regulations may be premature, local communities must increase their awareness of these impacts, and begin to incorporate these data in planning, design, and conservation issues.
7. Although the present study has shown that passive retreat is an important element of the shoreline response to anticipated global climate change, this inundation is certainly not the sole impact. Future research is mandated for other impacts on the coast of Massachusetts, including but not limited to:

- Effects of relative sea-level rise on groundwater resources.
- Effects of relative sea-level rise on marshes and other biotopes.
- Possible global climate change impact on storm climatology of Massachusetts waters.

ACKNOWLEDGEMENTS

This study has benefited from the contributions of many of our colleagues at the Woods Hole Oceanographic Institution. We are grateful for the assistance of Chris Pelloni of the Woods Hole office of the U.S. Geological Survey who provided coordinate translation of the numerical data. Jeff Benoit, of the Massachusetts Office of Coastal Zone Management, provided the foresight and management ability which made the study a reality.

REFERENCES


Figure 4
APPENDIX A

Hypsometry by Town:

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CHATHAM HYPSOMETRY
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Elevation (m.)

Percent of Upland Area

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A - 22
CHILMARK HYPSOMETRY
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Percent of Upland Area

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DANVERS HYPSOMETRY
CALCULATED FOR UPLAND 3m.+

Elevation (m.)

Percent of Upland Area
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CALCULATED FOR UPLAND 3m.+

Elevation (m.)

Percent of Upland Area

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CALCULATED FOR UPLAND 3m.+

Percent of Upland Area

Elevation (m.)

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3 10 20 30 40 50 60

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Elevation (m.)

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Elevation (m.)

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PLYMOUTH HYPSOMETRY
CALCULATED FOR UPLAND 3m.+

Elevation (m.)

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REHOBOOTH HYPSOMETRY
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Percent of Upland Area

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CALCULATED FOR UPLAND 3m.+
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Percent of Upland Area

Elevation (m.)

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0 10 20 30 40 50 60 70 80 90 100
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WAREHAM HYPSOMETRY
CALCULATED FOR UPLAND 3m.+

Elevation (m.)

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A-136
WESTPORT HYPSOMETRY
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YARMOUTH HYPSOMETRY
CALCULATED FOR UPLAND 3m.+
[PES.DIGDAT]
All digitizer output is stored here, as are the reformatted versions of the various data files. Reformatting can be executed here or in .HYP with inputs and outputs designated with [-.DIGDAT] preceding filename.

[PES.HYPDAT]
All individual town hypsometry data sets are stored here; all are named with a '.LIS' extension. These files also may be generated with an executable file stored in .HYP or one copied into this directory.

[PES.HYPDAT.TABLE]
All the tables associated with individual town hypsometry data sets are stored here. All are named with a '.TAB' extension.

[PES.HYPDAT.PLOTS]
All the plot files associated with the town hypsometry data sets are stored here. All are named with a '.PLT' extension.

B. PART ONE: STATE-WIDE HYPSOMETRY BY TOWN

1.) DATA BASE
A. The DEM data is in one-degree squares at 3 second intervals stored as south-to-north profiles. Elevations are to the nearest meter, formatted as follows:

* (8) 1024-byte virtual records making up a 8192-byte physical block. Only the first 1020 bytes of each virtual record contain data.

- The first block contains header information in the first and part of the second record
- Other Blocks:
  1st record-- 210 bytes (35 values in 6-byte fields) that are part of the previous profile
  2nd record-- 144 bytes of header for the next profile and 876 bytes (146 values) of data. The first profile of the data set starts at this point in the first block.
  3rd-8th record-- 1020 bytes (170 values) of data each. Each value is six bytes.
B. Land use, land cover, and political unit data sets cover the area of a USGS 1:250,000 scale map, and lie along west-to-east rows at 200 meter intervals on the UTM grid. The UTM grid is rotated counterclockwise just slightly with respect to longitude, yet bounded by the latitude boundaries of the USGS maps. Consequently some rows are shorter because they either run into the northern longitude bound or start along the southern longitude bound. These data are multiplexed and structured into 80-character records blocked into groups of 102 data points. The blocksize is 8160 bytes. The beginning and end of a block are unrelated to the beginning or end of a row. Each record contains the UTM coordinates to locate the point that it describes. This portion of the data base is hereafter referred to as the 'GRIDCELL' data as well.

2.) COORDINATE TRANSLATION

In order to locate the elevation data with respect to political units and land types (beaches, ponds, lakes, rivers being areas of special interest) the two data sets must be overlayed. To locate the land cover, etc. data with respect to the Lat./Long. system the United States Geological Survey (USGS) translated all the UTM coordinate pairs to Lat/Long values.

OUTPUT FORMATS USED BY USGS

For both the Boston and Providence data sets the translated Lat/Long coordinates that located each record were output separately (onto magnetic tape) maintaining the blocking factor of 102. We then merged (MERGEV3.FOR, MERGEV4A.FOR) these data with the appropriate fields from the original tape, creating 5100-byte blocks that contain 102 data points. Version 3 of MERGE handles the format used with the Providence sheet; version 4A handles the slightly different Boston sheet format. Each block of this new data set is made up of 102 50-byte virtual records. Each record contains ten bytes each for Latitude, Longitude, Land Cover Code, State/County Code, and Town (census) Code.

3.) REGRIDDING DATA

NOTE: GRDMATV3.FOR is for Providence data; GRDMATV4A.FOR is for Boston data. Input formats are the same, but data peculiarities necessitated customized error handling routines.

Once all the data describing each point in terms of land cover, political unit, and Lat/Long position were assembled on one tape, a new matrix was created that has data arranged on a grid identical to the DEM grid. GRDMATV3/V4A.FOR uses a simple method to create this DEM-based matrix of land and political codes. Essentially the software travels from point-to-point through an empty 3-second Lat/Long grid and searches for the closest data point in the UTM
C. PART TWO: HYPSOMETRY OF THREE HARBORS--
CONTOURING, COASTAL FLOODING, AND GROUNDWATER EFFECTS

NOTE: The full procedure for this operation is described only for Hyannis. Details that differ for the other harbors are outlined in following sections.

HYANNIS
1.) DATA GENERATION

Elevation Data: For this harbor a rectangle 4000 by 5000 meters was delineated, its southwest corner at UTM (391000,4609000). Within this area all contours were digitized, with different prefixes used to indicate the Z-coordinate or elevation values. Pond boundaries were input as separate contours. The shoreline was assigned an elevation of two feet relative to the NGVD of 1929 (this figure represents half the current tidal range plus six inches of sea level rise since 1929). Offshore contours were digitized, and contours valued at -3 and -4 ft. were added just offshore to help constrain the shoreline position. Approximately 11300 data points were recorded.

Scaling: A scale of 6.34969 meters per inch for 1:25,000 quadrangle maps yields output having a *10 scaling (units of decimeters). The origin must be entered as (0,0) so that the digitizer's five-digit integer output buffers are not overflowed. This offset is added back during reformatting in MINFORM.FOR and MINOFF.FOR, such that coordinate values are in the Universal Transverse Mercator System (UTM).

NOTE: A first attempt to generate this data set used a lower data density and thus produced only 6300 points (HY1*.DIG and .DAT). -30 ft. was used as the offshore contour depth, but was found to cause modelling problems onshore. Modelling the crucial low elevation areas (i.e., 1-10 ft.) presents a problem stemming from poor offshore (bathymetric) data. If only those offshore contours plotted on the topographic map are used to constrain the surface, shoals may be created just offshore and the shoreline is positioned poorly due to the extremely gentle slope in this area. Using artificially deep data points just offshore defines the position of the shoreline quite well, but will arch the onshore surface just inside the constraining shoreline data points. The compromise settled on is represented by the UPL4.DIG & DAT and OFF5.DIG & DAT combined into the HY5.DAT data file, and uses points of moderate depth to position the shoreline but not deform the onshore model to a great degree.
2.) FORMATTING DIGITIZER OUTPUT FOR MINCURV:

MINFORM.FOR and MINOFF.FOR are the programs used to put the digitizer output into X, Y, Z format for input to MINCURV.FOR. The digitizer will attach prefixes to its data output, but is limited to 14 different strings (AD-DD, OD-9D). To distinguish between all the different contours, pond boundaries, shoreline and shoreline structure vectors, as well as offshore contours, digitized data were processed in two batches where each prefix was translated into an appropriate elevation or depth. After formatting these two batches are merged into one file for gridding. Documentation contained in MINFORM.FOR and MINOFF.FOR explains how this was done for Hyannis. Other versions of these programs were written to accommodate other harbors. Perhaps a generic version will become appropriate that allows interactive assignment of certain Z-coordinate values to digitizer prefixes if many such maps are to be generated in the future.

3.) GRIDDING:

PUBLIC: MINCURV was used to generate an evenly spaced grid of elevation points. To retain full resolution within the single-precision storage arrays used by MINCURV, the first two digits of the northing and first digit of the easting value are not read in. This fix leaves the coordinates in a form that is still compatible with the USGS's UTM notations (i.e., our coordinates are what appear in upper case digits on the quadrangle maps). The Information Processing Center's (IPC) documentation of MINCURV is good. See our sample run as well.

A FEW HINTS:

- Remember to read data as F_.1 to account for the *10 scaling.
- Oversize the output grid to avoid loosing any data.
- When entering Min and Max values for X and Y, drop the appropriate leading digits.

MINCURV'S method: MINCURV performs a minimum curvature surface fit using the data provided. Therefore, slope breaks may affect the modelling of adjacent portions of the surface, especially in areas of low data density. This feature becomes troublesome when it becomes necessary to constrain the shoreline position more tightly than is done by the USGS 10-foot contours. To resolve the shoreline more closely, intuitively derived elevation data must be added immediately above and below the shoreline. A relatively steep nearshore slope is required to achieve the proper constraints, resulting in a slope break at the shoreline that can lead to an
Polygon plotting independent of the coastal contouring scheme was used to represent the enlargement of inland water bodies due to the rise in groundwater table. It was assumed that the groundwater table rise is equal to sea level rise.

'VECTOR.FOR' is a program written to create a '.POL' file interactively. It asks the user for certain data and reads in the plot data from a digitizer output file.

Creating the plot: To run HYPUC.FOR: Assign TT T41XX

HYPUC.FOR makes a UNIRAST.DAT file

To send plot to TK4695:
1) Assign the desired version of UNIRAST.DAT file to the most recent version number
2) Type 'ASSIGN TT T4695'
3) Type 'DUNIRAS TK4695'

To send plot to RASTEC:
1) Assign TT 41XX
2) Type 'DUNIRAS RASTEC/SC=SC'

WESTPORT

The Westport map area was established in a manner similar to the Hyannis area. Its origin is at UTM (324000,4696000) and is 4000 meters on a side. Only one inland water body is contained in the Westport map area, at one foot above mean high water. Flood level contours generated for the sea surface were considered legitimate for this coastal pond as well. Several versions of this data base exist; the latest is GL3.dat. At a scale of 1:25,000, there are 60.96012 meters per inch. Digitizer (*10) scaling is 6.0967012.

GLOUCESTER

Gloucester presented a radically different geology and land surface for flood modelling. Again the grid was set up along UTM lines with the origin at UTM (361500,4716500), extending 4000 meters east and north. Changes in the groundwater table were considered negligible here because of the relatively impermeable bedrock that contains it. Gloucester's steep and crenulated nearshore and its several extensive marshes made this harbor the most difficult to model. Four revisions were necessary. The current data base is GL4.dat. At a scale of 1:25,000, there are 60.96012 meters per inch. Digitizer (*10) scaling is 6.0967012.